







Methodological Considerations for the Use of Acid-Based Pre-Treatment Protocols for Carbon and Oxygen Analysis of Tooth Enamel

Karolina Varkulevičiūtė^{1,2} 🕞 | Christine Winter-Schuh^{1,2} | Cheryl A. Makarewicz^{1,2}

¹Institute of Prehistoric and Protohistoric Archaeology, University of Kiel, Kiel, Germany | ²Archaeology Stable Isotope Laboratory, University of Kiel, Kiel, Germany

Correspondence: Cheryl A. Makarewicz (c.makarewicz@ufg.uni-kiel.de)

Received: 26 August 2024 | Revised: 27 May 2025 | Accepted: 1 June 2025

Funding: This work was supported by the Deutsche Forschungsgemeinschaft, EXC 2150-390870439.

Keywords: carbon isotope | oxygen isotope | chemical pre-treatment | enamel diagenesis

ABSTRACT

Rationale: Chemical pre-treatment is a common methodological step aimed to remove exogenous materials introduced to archaeological tooth enamel in the burial environment through diagenetic processes. However, some of these methods, involving the use of oxidising reagents such as NaClO, H_2O_2 , as well as weak acids like CH_3COOH , have been shown to alter the chemical composition and stable isotope values of enamel. Here, we aim to re-examine the effects of commonly used pre-treatment protocols on bioapatite $\delta^{13}C$ and $\delta^{18}O$ values, and investigate the relationship between diagenetic alteration and measured isotope values, as indicated by pre-screening using attenuated total reflectance-Fourier transform infrared (ATR-FTIR) spectroscopy.

Methods: Modern and archaeological samples were subjected to 10 commonly used pre-treatment protocols that apply NaClO, $\rm H_2O_2$ and/or $\rm CH_3COOH$ to tooth enamel powders at treatment lengths. Preservation status and diagenetic alteration prior to and after treatment were investigated using ATR-FTIR. δ^{13} C and δ^{18} O values were measured before and after treatment to determine if different wet chemistry protocols induced isotopic shifts.

Results: The results show that all pre-treatment protocols imparted shifts in δ^{13} C and δ^{18} O values of up to $\pm 1.5\%$ in both archaeological and modern samples. Most treated samples display increased crystallinity, likely indicating sample recrystallisation. We suggest that these changes indicate the removal of contamination and diagenetic alteration, and also the dissolution and restructuring of enamel carbonate leading to changes in the in vivo isotope signal.

Conclusions: We discourage the use of H_2O_2 and NaClO to remove organic matter from samples as it incurs unwanted changes to the enamel structure and carbon and oxygen isotope ratios. We also recommend the use of only short-duration acetic acid treatment protocols to avoid recrystallisation caused by prolonged acid exposure and concomitant unwanted change to in vivo isotope values.

1 | Introduction

Carbon (δ^{13} C) and oxygen (δ^{18} O) isotope ratios recorded in tooth enamel are routinely measured by ecologists and archaeologists

to establish (paleo)environment and vegetation dynamics as well as dietary intake and mobility in humans and animals [1–5]. For example, questions regarding the introduction of new plant cultivars, such as maize and millet, can be tracked in human and

This is an open access article under the terms of the Creative Commons Attribution License, which permits use, distribution and reproduction in any medium, provided the original work is properly cited.

© 2025 The Author(s). Rapid Communications in Mass Spectrometry published by John Wiley & Sons Ltd.

animal teeth carbon isotope values by taking advantage of the strong isotopic distinction between C_3 and C_4 plants [6], while oxygen isotope values have been used to investigate human mobility as it directly relates to $\delta^{18}\mathrm{O}$ values of imbibed water of their local residence during the tooth formation process [5]. Oxygen isotope ratios recorded in animal teeth have high potential to generate high-resolution paleoclimate information at local scales that can be more directly linked to loci of ancient human activity and also provide information about animal mobility.

The accuracy of such dietary, mobility, and palaeoclimate reconstructions, however, is highly dependent on the recovery of in vivo isotope signals from tooth enamel bioapatite. Bioapatite can be altered by several modes of diagenetic alteration. This includes the exogenous introduction of contaminating organic and inorganic compounds [7], enamel structure dissolution [8], and recrystallisation [9-11] that can take place after teeth enter the archaeological burial environment and, consequently, alter δ^{13} C and δ^{18} O values. Chemical pre-treatment of samples, while intended to eliminate or reduce the impact of diagenetic alteration on isotope ratios, can introduce additional changes to the enamel bioapatite and thus change its isotopic composition [12-15]. Whether caused by diagenetic alteration or pre-treatment chemistry, shifts in oxygen and carbon isotope ratios away from the original in vivo values can profoundly affect interpretation of results. For example, alterations to in vivo carbon isotope values may result in inaccurate estimation of the contribution of C₄ plants such as millet to the diet and this, in turn, can lead to differing interpretations of the importance of this cultivar in human subsistence or livestock foddering regimes. Similarly, shifts in oxygen isotope ratios can lead to different conclusions about mobility patterns, seasonality, and local environment.

1.1 | Tooth Enamel Composition and Diagenesis

Tooth enamel is comprised of a mineral component (>96%) mostly in the form of hydroxyapatite (Ca₁₀(PO₄)₆(OH)₂) and < 1% organic material [16]. Minor elements, such as magnesium (Mg), sodium (Na), chlorine (Cl), and strontium (Sr), are also commonly present, filling vacancies at the calcium (Ca) and hydroxide ion (OH⁻) sites of the hydroxyapatite crystalline lattice [17, 18]. The inorganic phase of tooth enamel consists of flat, hexagonal hydroxyapatite crystallites [19] arranged as micrometresized prisms (or rods) filling the enamel matrix [20]. Within the prism 'body', crystallites are densely packed with little interstitial pore space. However, the volume of pore space increases at the 'tail' region of the prism and along prism boundaries due to changes in crystal orientation [21, 22]. Porosity within enamel prism structures facilitates the diffusion of water [23] and watersoluble matter throughout the tooth enamel, including ions and some small organic molecules [24]. Circulation of water in the pore space of bioapatite can cause dissolution of organic and inorganic matter as well as introduce exogenous compounds via ion exchange [25], potentially altering the original crystalline structure of the enamel.

Diagenetic alteration of the bioapatite structure occurs via incorporation of new chemical species, such as exogenous carbonate or minor elements [26–28], dissolution of structural carbonate [8], or internal restructuring of the bioapatite [29]. Incorporation

of non-biogenic carbonates may occur in carbonate-rich burial environments where exogenous carbonates precipitate in enamel cracks or as bicarbonate adsorbed to crystal surfaces [26]. Biological molecule degradation initiated post-mortem increases the number and size of interstitial pores present in the enamel prism structure, which allows fluid from the surrounding environment to penetrate and migrate through the enamel matrix [30]. Consequently, minor elements present in soil burial environments such as magnesium (Mg), iron (Fe), strontium (Sr), and zinc (Zn), diffuse through liquid along prism boundaries and are subsequently incorporated into the extant bioapatite structure [9, 31, 32]. This diffusion of trace elements, such as strontium, into enamel bioapatite has been associated with shifts in the isotopic composition of tooth enamel [33–35].

However, the impact of minor element leaching on the enamel matrix remains unclear. Enamel crystallinity may be affected and, consequently, carbon and oxygen isotope values. Dissolution and recrystallisation of bioapatite, initiated by changes in burial environment temperature, pH, and microbial action, may shift in vivo stable isotope values [8, 36, 37]. Heating bioapatite removes structurally incorporated water and organic components [38]. Alkaline conditions can promote recrystallisation of bone bioapatite [39]. Meanwhile, microbial activity can break down molecular bonds within the bioapatite, thus accelerating isotope exchange between enamel carbonate and phosphate oxygen and groundwater [40].

The effect of diagenesis on bone and tooth bioapatite can be identified and quantified using attenuated total reflectance-Fourier transform infrared (ATR-FTIR) spectroscopy. Certain characteristic FTIR peaks can indicate the presence of exogenous material. For example, a peak at approximately 710 cm⁻¹ is diagnostic of calcite (CaCO₃), which can occur as a precipitate from soil [41, 42]. FTIR indices variously describe the ordering of bioapatite crystalline structure and carbonate substitutions. The infrared splitting factor (IRSF) describes bioapatite crystallinity and can be useful in identifying heating or recrystallisation. Unaltered enamel typically shows IRSF values between 3.1 and 3.9, while values above 4.0 indicate heating [43, 44] or recrystallisation caused by exposure to groundwater [10, 11, 45-47]. Other common indices include ratios of A-type and B-type carbonate substitutions (type A/type B carbonate to phosphate index, API and BPI) and carbonate to phosphate ratios (C/P). These indices, however, and the diagenetic alteration communicated by these indices may not necessarily mean that concurrent shifts in the isotopic composition of bioapatite from their original in vivo δ^{13} C and δ^{18} O values took place.

1.2 | Common Methods to Address Diagenesis

Concern that diagenetic alteration may alter the isotopic composition of the target tissue has precipitated the development and use of various chemical pre-treatments, each aimed at removing endogenous organic matter, organic contamination, and/or exogenous diagenetic carbonates with the goal of measuring carbon and oxygen isotope ratios that most faithfully reflect original organismal δ^{13} C and δ^{18} O values (summary in Table 1). Oxidising reagents such as basic sodium hypochlorite (NaClO) are used to remove organic matter in concentrations ranging from 1% to

TABLE 1 | A summary of enamel pre-treatment protocols commonly applied to tooth enamel powder.

References	[48–50]	[14, 51–55]	[15, 56–59]
Concerns	Might not fully remove organic contamination, induces absorption of atmospheric CO_2	Might not fully remove organic contamination, low pH causes excessive sample loss	Removal of structural carbonate, excessive sample loss, recrystallisation
Treatment length	30 min-24 h	15 min-24 h	10 min-24 h
Temperature	Room temp.	Room temp. or heated to 70°C	Room temp.
Concentration	1%-5%	3%-30%	0.1–1 M
Purpose	Removal of organic content	Removal of organic content	Removal of carbonate contamination
Reagent	NaClO	$\mathrm{H_2O_2}$	СН3СООН

5% at treatment lengths up to 24h at room temperature [48–50]. Reactions of bioapatite with NaClO also appear to cause a dissolution of atmospheric CO2, which is converted into carbonic acid followed by bicarbonate and carbonate ions precipitating on the powder sample, thus altering the carbon and oxygen isotope ratios from their original isotope values [12]. A less prevalent, but still popular, oxidising reagent is acidic hydrogen peroxide (H₂O₂), applied at 3% to 30% strength with treatment lengths varying from 15 min up to 24h at different prescribed temperatures [51–55]. The low pH of H₂O₂ causes excessive sample loss [14], problematic for high-resolution sampling strategies that generate small amounts of material for isotope analysis such as intra-tooth sequential sampling. The loss of structural carbonate post H₂O₂ treatment is also reflected in the decrease of type B carbonate [60, 61], a carbonate replacing a phosphate in the hydroxyapatite structure. Despite these concerns and the negligible organic content in enamel, both reagents are still widely applied to ancient specimens.

Exogenous carbonates from tooth enamel powders are usually removed using acetic acid (CH2COOH), following foundational experimental work by Lee-Thorp and van der Merwe [56]. Typically, non-lattice-bound and therefore more soluble, exogenous carbonates are easily removed with low-concentration acids. In recent years, 0.1 M acetic acid and 1 M sodium acetate buffered acetic acid applied at room temperature with treatment length varying from 10 min to 24h are most commonly used [15, 57-59], contrasting with the use of 1M concentration unbuffered acetic acid in the past [62, 63]. Buffered acetic acid is characterised by higher, closer to neutral, pH levels that help conserve sample amount [60]. However, assessing the effectiveness of acetic acid in removing exogenous carbonates is challenging, as these carbonates share a chemical composition and therefore FTIR absorption peaks, with indigenous carbonates. Only the significant presence of calcite, which displays a characteristic peak at 710 cm⁻¹ not present in pure enamel hydroxyapatite [41, 42], or extensive recrystallisation of carbonate into brushite [56] can definitively identify diagenetic alteration of carbonate. Furthermore, it appears that structural lattice-bound carbonates are also lost during pre-treatment, suggested by sample loss and decreased carbonate concentrations observed in powdered samples from modern specimens that never entered a burial environment [12, 14, 60]. Prolonged acetic acid treatment time can also lead to the recrystallisation of carbonate recrystallisation into brushite [56], although it is likely that some degree of recrystallisation takes place even with shorter treatment durations [13]. Acetic acid treatments have been shown to inconsistently impact δ^{13} C and δ^{18} O values, with shifts up to approximately $\pm 1\%$ observed [12–14]. Notably, it is not well understood whether these isotopic shifts are due to the removal of contaminating exogenous carbonates, removal of biogenic carbonates, enamel recrystallisation, or introduction of exogenous matter [12, 14, 15, 59, 60].

As archaeological interpretations of paleoclimate and paleodiet increasingly draw from ever smaller per mil isotopic distinctions between samples to identify fine-scale changes in environment, mobility, and consumption behavior, obtaining isotope values that most closely reflect their in vivo signal is increasingly important. Here, we present a large scale, systematic re-examination of the extent to which different pre-treatment protocols affect the structural integrity, chemical composition, and isotopic composition of tooth enamel bioapatite. We aim to illustrate how different pre-treatment protocols alter enamel bioapatite and how pre-treatment led changes may affect measured isotope values, and, thereby influence subsequent palaeodietary, palaeomobility, and palaeoclimatic interpretations.

2 | Materials

To assess the impact of different pre-treatment protocols, a total of 30 archaeological and 10 modern sheep/goat (Ovis aries/Capra hircus) and cattle (Bos taurus) mandibular molar teeth (M/2 and M/3) were sampled. We have chosen to investigate sheep/goat and cattle as they are two of the most commonly found species in the archaeological record in Eurasia as well as most commonly investigated taxa in enamel stable isotope studies. Archaeological sites of different ages, geographic, and climatic contexts were chosen to explore how these factors might impact enamel bioapatite response to pre-treatment. Based on visual inspection, selected specimens displayed varying states of preservation. In addition to archaeological specimens, modern faunal material was chosen as a reference against which diagenetically altered material could be evaluated. Archaeological material was obtained from three sites: Tell Handagug South (Jordan), Wadi Faynan 16 (WF16, Jordan), and Oldenburg LA 77 (Germany). Modern samples were selected from two sites: Olfen (Germany) and Baga Gazaryn Chuluu (Mongolia).

Tell Handaquq South (THS; n=10; five Ovis/Capra, five Bos) is a large Early Bronze Age settlement site (3600–2000 cal BC) located on a slope abutting the southern bank of the Wadi Zarqa at the base of the Jordanian highland escarpment [64]. Regional paleoclimate studies suggest the site experienced semi-arid conditions with estimated precipitation levels of approximately 350 mm per annum [65, 66]. Calcic soils characterise the area around the settlement [67], and while no sediment analyses have been carried out at the site, most of the deposits are described as silty in the excavation report [68]. The majority of the bones and teeth recovered from Tell Handaquq South were encrusted with calcareous concretions of up to a few millimetres in thickness and required mechanical removal prior to the tooth sampling procedure (Figure S1A).

Oldenburg LA 77 (OLD; n=10; four Ovis/Capra, six Bos) is a large Middle Neolithic Funnel Beaker settlement site (ca. 3300-2800 BC) located on the Baltic coast on the Wagrien peninsula in Schleswig-Holstein, northern Germany [69]. The site is situated on the waterfront wetland area within the Oldenburger Graben glacial valley and, during occupation two former fjord-like estuaries (Oldenburg and the Gruber Förde) formed lagoons with varying water table and salinity levels [70, 71]. The climate in the region during the Neolithic was characterised by cool temperatures and wet conditions [72, 73]. The peripheral areas of the site support peat formations, calcareous, detritus-rich mud with sand inclusions [71]. ATR-FTIR analysis of the sediment indicates a high silica content (likely quartz), gypsum, and only very low quantities of carbonates (Figure S4A). Oldenburg LA 77 specimens sampled here display brown and, sometimes, almost black staining (Figure S1B). Such staining on bone and teeth, common for specimens recovered from waterlogged contexts, is

likely due to contact with dark organic matter, iron, and manganese compounds found in soil [74]; high concentrations of iron and manganese oxides are often found in such teeth [75–77].

Wadi Faynan 16 (WF16; n = 10; five Capra sp., five Bos primigenius) is a Pre-Pottery Neolithic A settlement (10070-8160 cal BC) [78, 79] located at the head of the Wadi Faynan, a seasonally flooded water course between the rocky Sharah mountains rising up to the Jordanian plateau (ca. 1300 m a.s.l.) and the arid flats of Wadi Araba (ca. 0 to 200 m a.s.l.). Hydrological simulations suggest that at the beginning of the Holocene (ca. 12000 BC), Wadi Faynan experienced similar, albeit slightly wetter, conditions than the present-day climate, characterised by rapid flood responses to rainfall events in the winter season [80]. Sedimentology work shows that sediments recovered from the archaeological deposits were sandy, alkaline, and primarily minerogenic in nature, characterised by varying proportions of quartz, calcite, gypsum, and a low organic content [81]. This is further confirmed by FTIR analysis of sediment, recovered with the bone specimens, that demonstrate a high silica and carbonate content with the presence of quartz and gypsum (Figure S4B). The animal teeth recovered from WF16 present as highly fragmented with thin and brittle enamel (Figure S1C).

Modern sheep teeth (n=5) were chosen from an ethnographic site located in the Mongolian Gobi steppe desert at Baga Gazaryn Chuluu (BGC; Dundgovi aimag). Animals died in 2000-2001 during a dzud event, an extremely cold winter accompanied by snow or ice that removes or covers graze so that livestock starve. Summers are hot, with most precipitation falling during the summer months, and winters severely cold [82]. Specimens were exposed until collection in 2003 and bore signs of weathering. Modern cattle teeth (n=5) were sourced from Olfen (OLF) located in North Rhine-Westphalia, western Germany, from animals slaughtered in 2018-2019. The cattle were largely unmanaged and kept in a grassy pasture field year-round with some hay supplementation during the winter months. The region is characterised by a mild temperate climate, with an average annual temperature of 10° C ($\approx 18^{\circ}$ C in July and $\approx 5^{\circ}$ C in January) and mean annual precipitation of 850 mm.

3 | Methods

Prior to sampling, the tooth enamel surface was cleaned from soil and concretions using a dental burr, followed by ultrasonication in ultrapure water and then dried. This was particularly important for teeth from THS which were covered in thick calcareous concretions; removed concretions were set aside for later ATR-FTIR analysis. The length of the tooth lobe was cut, and dentine was removed using a diamondtipped Dremel drill. Dentine is characterised by a different colour and hardness compared with enamel, and full removal of the dentine was ensured through visual inspection under good lighting conditions. Clean enamel was ground into a homogenous powder using a Retsch MM40 mixer mill. Prior to experimental treatment, each enamel sample was screened using ATR-FTIR spectroscopy on a Bruker Alpha II ATR-FTIR at the Archaeology Stable Isotope Laboratory, Institute of Prehistoric and Protohistoric Archaeology, University of Kiel. Around 1 mg of sample (enough to cover the crystal) was

measured twice, and the spectra from 450 to 4000 cm⁻¹ were collected for 100 scans at a resolution of 4 cm⁻¹. Baseline correction and peak picking were carried out using the Bruker OPUS 8.7 software. The following indices were calculated to characterise the composition and structure of the enamel and assess any diagenetic alteration present in the samples: infrared splitting factor (IRSF), ratios of A-type (substituting the OH group), and B-type (substituting the phosphate) carbonate to phosphate peaks (API and BPI), carbonate to phosphate ratio (C/P), and carbonate to carbonate ratio (C/C) (Table 2). Guidelines for accepted ranges for 'well-preserved' enamel were taken from France et al. [83], which combined the observations of modern and historic samples that passed visual analyses, weathering indices, and organic collagen preservation indicators. While collagen preservation does not equate to the preservation of the in vivo carbonate-derived isotope values, it is still a useful control measure in the absence of other carbonate preservation indicators [83].

In this study, 11 different treatment protocols were tested, including an untreated control group. Protocols included organic removal pre-treatment with $\rm H_2O_2$ and NaClO as well as secondary carbonate removal pre-treatment with different concentrations of acetic acid (Table 3). For each treatment, 20 mg of powder was subsampled and treated with around 1.5 mL of solution, and 15 mg of enamel powder was subsampled for the untreated group.

Three organic removal treatment protocols were chosen for this study based on the most commonly cited protocols in the literature. Group 1 received a 3% $\rm H_2O_2$ pre-treatment applied at room temperature for 15 min. Groups 2 and 3 received a 3% NaClO pre-treatment at room temperature for 45 min and 16 h, respectively. Treatment lengths were selected to capture the large range in

TABLE 2 | ATR-FTIR indices used in this study, along with corresponding peak height ratios, significance, and the suggested ratios for "well-preserved" enamel (based on France et al. [83]).

Name	Peak height ratios (cm ⁻¹)	Significance	"Well- preserved" range
IRSF	(565+605)/590	Indicator of crystallinity	3.1–3.9 (untreated) 3.1–4.0 (chemically treated)
C/C	1455/1415	Compares two carbonate bands	0.9-1.1
C/P	1415/1035	Carbonate to phosphate ratio	0.08-0.2
API	1540/605	Amount of A type carbonate substitutions	0.04-0.2
BPI	1415/605	Amount of B type carbonate substitutions	0.1-0.65

duration of treatment, from 30 min to several days, described in the recent literature [86, 88, 89].

Secondary carbonate removal was tested with 0.1 M unbuffered acetic acid ($\mathrm{CH_3COOH}$) and 1 M calcium acetate buffered acetic acid (($\mathrm{CH_3COOH}$)₂Ca), two commonly used reagents used for this purpose [15, 56–59]. These two types of acetic acids were also chosen to try and replicate the differences in mass loss resulting from pre-treatment previously investigated by Snoeck and Pellegrini [60]. Three treatment durations were selected for both types of acetic acid protocols: 15 min, 4h and 8 h.

The last treatment protocol used here is a combined protocol of 45 min 3% NaClO treatment followed by a 15-min 1M buffered acetic acid treatment (i.e., [49, 86, 90–92]). The 15-min buffered acetic acid treatment has been chosen for the second step as previous experimental work has shown that buffered acetic acid appears to reduce sample loss compared to the unbuffered acetic acid treatments [12, 13, 60].

After pre-treatment, all of the sample powders were rinsed with ultra-pure water five times and freeze-dried prior to isotope ratio mass spectrometry. Mass loss caused by pre-treatment was determined by weighing each sample prior to and after the treatment. Supernatant remaining after the treatment and rinsing was dried and analysed using a Bruker Alpha II ATR-FTIR at the Archaeology Stable Isotope Laboratory, Institute of Prehistoric and Protohistoric Archaeology, University of Kiel.

Stable oxygen and carbon isotope ratios of powdered enamel carbonates were analysed using a MAT 253 dual-inlet isotope ratio mass spectrometer coupled to a Kiel IV carbonate device at the Leibniz-Laboratory for Radiometric Dating and Stable Isotope Research, University of Kiel. Samples were reacted by individual acid addition (99% $\rm H_3PO_4$ at 75°C), and the evolved carbon dioxide was analysed eight times for each sample. Isotope ratios are reported in permil (‰) using the delta (δ) notation and calibrated relative to the Vienna PeeDee Belemnite (VPDB) scale. The δ notation refers to a measure of deviation in ratio between a heavy and a light stable isotope (in this case $^{18}\rm O/^{16}\rm O$ and $^{13}\rm C/^{12}\rm C$), and the result expresses the relative difference between a sample and a standard in parts per thousand (‰).

$$\delta^{18}O = \left(\frac{\left(\frac{^{18}O}{^{16}O}\right)\text{sample}}{\left(\frac{^{18}O}{^{16}O}\right)\text{standard}} - 1\right) \times 1000$$

$$\delta^{13}C = \left(\frac{\left(\frac{^{13}C}{^{12}C}\right)\text{sample}}{\left(\frac{^{13}C}{^{12}C}\right)\text{standard}} - 1\right) \times 1000$$

Analytical precision was $\pm 0.08\%$ for both $\delta^{18}O$ and $\delta^{13}C$, based on the performance of international standards (NBS19: $\delta^{13}C+1.95\%$, $\delta^{18}O-2.20\%$; IAEA-603: $\delta^{13}C+2.46\%$, $\delta^{18}O-2.37\%$) and internal laboratory pure carbonate standards (Hela $\delta^{13}C+0.91\%$, $\delta^{18}O+2.48\%$; SHK: $\delta^{13}C+1.74\%$,

10970231, 2025, 20, Downloaded from https://analyticalsciencejournals.onlinelibrary.wiley.com/doi/10.1002/cm.10090 by Christian-Albrechts-Universitat zu Kiel, Wiley Online Library on [22/10/2025]. See the Terms and Conditions (https://onlinelibrary.

ols inwestigated in this study	dis mivestigated in tims study.
Dre-treatment protoco	TIC-II CHITICIII DI OLOCI
TARIE 2	CHICAL

Lahel	Reagent	Concentration	Time	T (°C)	Instification	Previous annlications
	14					J.J.
	No treatment					
$15\mathrm{min}~\mathrm{H_2O_2}$	$\mathrm{H_2O_2}$	3%	15 min	Room temp.	Oxidising agent used for organic contamination removal (e.g., humic acids) in bone mineral, subsequently used for enamel preparation	[51–53, 84]
45 min NaClO	NaClO	3%	45 min	Room temp.	Oxidising agent used for organic contamination	[49, 85–87]
16 h NaClO			16 h		removal (e.g., humic acids) in bone mineral, subsequently used for enamel preparation	
15 min unbuff	CH_3COOH	$0.1\mathrm{M}$	15 min	Room temp.	Exogenous carbonates are more soluble than	[49, 56, 85, 86]
4h unbuff			4 h		apatite and can be removed by leaching in dilute acid such as CH COOH	
8h unbuff			8 h			
15 min buff	$(CH_3COOH)_2Ca$	1 M buffered	15 min	Room temp.	Buffering the acid makes it closer to neutral	[12, 60, 86]
4h buff			4 h		pH and less destructive to the sample	
8h buff			8 h			
NaClO + Acetic	NaClO+CH ₃ COOH	$3\% + 1 \mathrm{M}$ buffered	45 min + 15 min	Room temp.	Combining the organic removal step and carbonate exogenous carbonate removal protocols gets rid of all unwanted matter	[49, 85–87]
					F 0 1	

 $\delta^{18}{\rm O}$ –4.85‰; HB1: $\delta^{13}{\rm C}$ –12.10‰, $\delta^{18}{\rm O}$ –18.10‰). Additional laboratory internal enamel bioapatite standards (ER1: $\delta^{13}{\rm C}$ –12.18‰, $\delta^{18}{\rm O}$ –7.78‰; ES1: $\delta^{13}{\rm C}$ –13.76‰, $\delta^{18}{\rm O}$ –6.41‰) yielded a standard deviation of 0.07‰ for $\delta^{13}{\rm C}$ and 0.10‰ for $\delta^{18}{\rm O}$ (ER1), and 0.07‰ for both $\delta^{13}{\rm C}$ and $\delta^{18}{\rm O}$ (ES1). Sample duplicates were run on average every 10 samples (n=84), showing an average deviation of $\pm 0.05\%$ for $\delta^{13}{\rm C}$ and $\pm 0.16\%$ for $\delta^{18}{\rm O}$.

Concretions removed from THS teeth were ground into a homogenous powder. A subsample of the powder was analysed using ATR-FTIR spectroscopy and isotope ratio mass spectrometry to determine the chemical composition and well as δ^{13} C and δ^{18} O values of the concretions. Sediments adhering to the Oldenburg LA 77 and WF16 teeth or sediments remaining in the sample bags were also selected and analysed using ATR-FTIR spectroscopy to determine the chemical composition of the burial environment.

4 | Results

4.1 | Mass Loss

Sample loss due to treatment protocol type is expressed here as a percentage loss of dried sample calculated relative to the sample weight prior to treatment (see Figure 1). All treatment protocols caused a notable mass loss in both archaeological and modern samples ranging on average from 16.3% (3% NaClO 45 min) to 24.6% (4h unbuffered CH₂COOH). There was no correlation observed between the treatment length and sample loss in both buffered and unbuffered acetic acid treatments (r=-0.13 and r=0.084, respectively). A nonparametric Kruskal-Wallis test followed by a post hoc Dunn's test (see Table S5) showed that only samples treated with 3% NaClO for 45 min were statistically different from the other treatment groups with an average mass loss of 18.3% compared with a mass loss of 22-24% observed in other groups. Outliers are present in all of the treatment groups; the higher sample loss in the outliers are likely due to static in the sample tubes during freeze drying, storage, and weighing.

4.2 | Sample Preservation Parameters Prior to Treatment

Sample preservation was assessed through the calculation of five commonly used FTIR indices: IRSF, C/C, C/P, API, and BPI, and the ranges for 'well-preserved' enamel published in France et al. [83] (see Table 2). The majority of the samples were within the accepted range for all indices (Figure 2; Table S2). The API index, which represents ratios of A-type carbonate substitutions to phosphate peaks, was most frequently expressed outside of accepted range, visible in nine out of 40 samples. This included two modern specimens from BGC which was unexpected, as modern samples should not be diagenetically altered. The C/C index, the ratio between type A and type B carbonates, yielded values above the accepted range of 1.1 for four out of 40 samples. Values outside of the accepted range for C/P, the ratio between carbonate and phosphate peaks and IRSF, crystallinity factor, were observed in two out of 40 samples. For samples THS2 and OLD5, IRSF values were 4.0 and 4.3, respectively, considerably above the accepted 3.1-3.9 range for untreated enamel and much higher compared to all other specimens, which on average yielded values of 3.5. The higher IRSF values indicate a more ordered crystal structure within enamel, which could be a result of carbonate alteration, ionic exchange during burial, microbial activity, or recrystallisation caused by heating or natural recrystallisation during burial/groundwater exposure [10, 11, 43-47].

Modern samples from OLF and BGC yielded acceptable preservation parameters. Two samples from BGC failed the API index but this could be related to visible weathering in the specimens or the occasionally indistinct peak at 1540 cm⁻¹ peak used in the calculations of this parameter. All of the samples from WF16 yielded preservation parameters within the accepted range for all five indices, suggesting good preservation of the enamel despite outstandingly poor dentine collagen preservation. Samples from THS and Oldenburg LA 77 yielded the highest proportion of samples, seven and five, respectively, outside of accepted ranges. The 1415 cm⁻¹ peak used in the calculation of C/C, C/P and BPI indices, all of which yielded values outside accepted ranges at THS and Oldenburg LA 77,

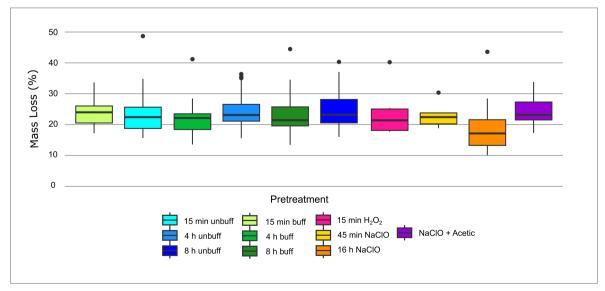


FIGURE 1 | Sample weight loss after treatment according to pre-treatment protocol and treatment duration.

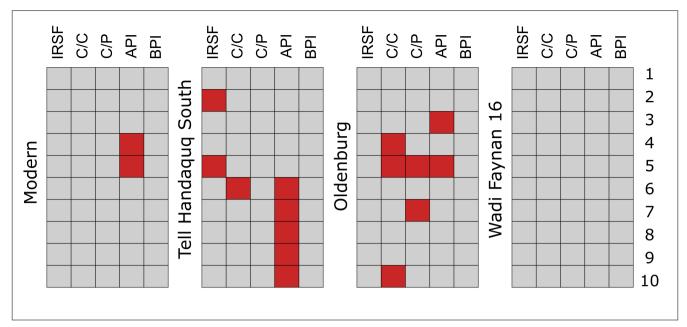


FIGURE 2 | Summary of FTIR parameter pre-screening results. Each box represents a powdered enamel sample from an individual tooth specimen (numbers 1 through 10) according to site: Modern (BGC and Olfen), Tell Handaquq South, Oldenburg LA 77, and Wadi Faynan 16. Grey boxes indicate that the powdered sample yielded peak height ratios within the accepted range for well-preserved enamel; red boxes indicate the sample yielded peak height ratios outside the accepted range for specific parameters presented in Table 2.

is a prominent peak also observed in pure $CaCO_3$, $MgCO_3$, and other carbonates commonly found in soil [83]. The $1455\,\mathrm{cm^{-1}}$ peak, used in the calculations of C/C index, often overlaps with secondary (diagenetic) carbonates [83]. This coincides well with the visual observations of samples from these sites. Calcareous concretions covering the THS specimens may have promoted precipitation of exogenous carbonate into the enamel structure. Brown discolouration of the Oldenburg LA 77 teeth, commonly observed in specimens from waterlogged contexts, likely results from the absorption of manganese and iron oxides from the surrounding soil [74–77], and may also signal post-depositional diagenesis.

4.3 | Effects of Pre-Treatment on the δ^{13} C and δ^{18} O Values of Enamel

To assess the isotopic shifts induced by different treatment protocols, isotope values measured from pre-treated samples were subtracted from those of untreated samples ($\Delta_{\rm isotope} = \delta_{\rm treated} - \delta_{\rm untreated}$). Positive values indicate an increase in $\delta^{\rm 13}{\rm C}$ or $\delta^{\rm 18}{\rm O}$ values, while negative values correspond to a decrease in $\delta^{\rm 13}{\rm C}$ or $\delta^{\rm 18}{\rm O}$ values (Figure 3; Table S1).

4.3.1 | Effects of Organic Removal Pre-Treatments on Isotope Composition

For $\rm H_2O_2$ treated samples, virtually no carbon isotopic shift was observed, with $\Delta^{13}C$ values ranging from -0.2% to +0.1%. NaClO, on the other hand, had a stronger and more variable effect on carbon isotope values. Oldenburg LA 77 $\Delta^{13}C$ values were positive post-treatment (range $\Delta^{13}C = +0.1\%$ to +0.6%), while WF16 $\Delta^{13}C$ values were mostly negative (range $\Delta^{13}C = -0.4$ to 0.0%). For THS and modern samples, isotopic shifts in

both directions were observed (range $\Delta^{13}C=-0.2\%$ to +0.4% and -0.3% to +0.4%, respectively). Sample powders subjected to both a 45 min and a 16 h NaClO pre-treatment exhibited slightly larger shifts in $\Delta^{13}C$ values for the 16 h treatment compared to the 45 min treatment, with differences of up to 3‰. All organic removal pre-treatment protocols imparted carbon isotope shifts of up to 0.6% and there were no predictable patterns in the range or directionality of change in $\Delta^{13}C$ values shared between samples from specimens from different sites.

 $\Delta^{18}{\rm O}$ values were mostly negative for samples treated with ${\rm H_2O_2}$ (range $\Delta^{18}{\rm O}=-0.4\%$ to 0.1%) and 16 h NaClO (-0.9% to 0.0%) across all sites. The 45 min NaClO treatment resulted in both positive and negative $\Delta^{18}{\rm O}$ values ranging from -0.4% to +0.2% at THS, -0.2% to +0.1% at Oldenburg LA 77, -0.9% to -0.2% at WF16, and -0.2% to +0.4% for modern samples. When the shorter 45 min and longer 16 h NaClO treatment was applied to the same enamel powder, $\Delta^{18}{\rm O}$ values were up to 0.6% larger for the 16 h treatment compared to the 45 min treatment. For samples treated with NaClO, the shift in $\Delta^{18}{\rm O}$ values was minimal in most modern, THS, and Oldenburg LA 77 samples. In sharp contrast, samples from WF16 were strongly affected by the oxidising organic removal pre-treatment protocol, with a shift of -0.8% or more visible in six out of 10 samples.

4.3.2 | Effects of Secondary Carbonate Removal Pre-Treatments on Isotope Composition

The change in carbon isotope ratios observed post-acetic acid treatment for specimens from all sites resulted in negative $\Delta^{13}C$ values, with the exception of Oldenburg LA 77 where the direction of isotopic change varied between individual tooth specimens. Shifts in $\Delta^{13}C$ values between 0.1 M unbuffered and 1 M buffered acetic acid solutions was minimal, in general below

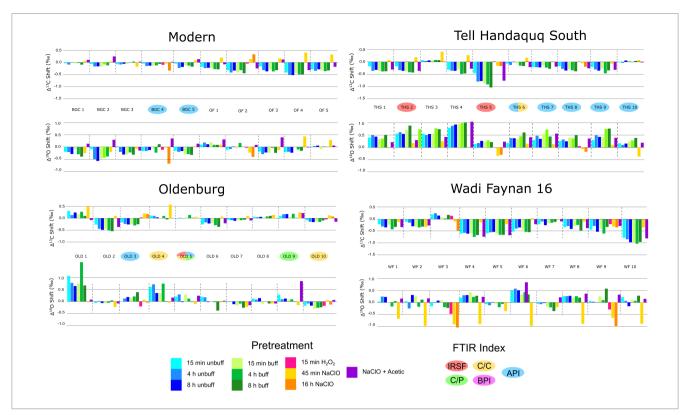


FIGURE 3 | Shifts in Δ^{13} C and Δ^{18} O values after application of different experimental pre-treatment protocols. An oval laid over the sample name indicates the FTIR parameters outside the accepted range for "well-preserved" enamel prior to treatment for that sample.

 $\pm 0.3\%$. There was no observable correlation between treatment length and the amplitude of isotopic change in both the buffered and unbuffered acetic acid treated groups. The shifts in $\Delta^{13} C$ values observed in the Oldenburg LA 77 and BGC modern specimens were smaller and generally negligible compared with the other sites. For specimens from THS, WF16, and Olfen, shifts in $\Delta^{13} C$ values generally ranged from -0.2% to -0.5%, with a few isotopic outliers visible across the specimens.

The directionality of oxygen isotopic change visible in samples varied by site, with a positive shift visible in $\Delta^{18}O$ values (+0.1% to +1.0%) at THS, a negative shift (-0.6% to +0.2%) observed in modern specimens, and specimens from both Oldenburg LA 77 (-0.4% to +1.6%) and WF16 (-0.3% to +0.5%) exhibiting both negative and positive shifts in $\Delta^{18}O$ values. While there was a wide range in the amplitude of oxygen isotopic change visible for samples treated with buffered and unbuffered acetic acid (ranging from $\pm 0.0\%$ to $\pm 1\%$), the majority of the samples, regardless of treatment length or type of acetic acid used, exhibited minimal isotopic differences $\pm 0.5\%$ or lower. Only samples from THS exhibited more pronounced oxygen isotopic differences between buffered and unbuffered acetic acid treatments.

4.3.3 | Effect of Combined Pre-Treatment on Isotope Composition

The direction and magnitude of carbon and oxygen isotopic change in samples treated with combined NaClO and acetic acid

protocols are similar to those observed for acetic acid-only treatments. However, combined treatment shifts $\Delta^{13}C$ and $\Delta^{18}O$ values in a positive direction while the NaClO-only treatment shifts $\Delta^{13}C$ and $\Delta^{18}O$ values in a negative direction. Isotopic shifts in the combined treatment group vary widely across individual specimens (range = -0.8% to +1.0%), and, no site-specific patterns are visible.

4.3.4 | Effects of Pre-Treatment on Enamel ATR-FTIR Spectra

The effect of different treatments on ATR-FTIR spectra is visible in a few of the samples. Samples from Oldenburg LA 77 and WF16 yielded distinctly distorted spectra after pre-treatment (Figure 4). Most samples also failed one or more FTIR parameters post-treatment, even those that initially yielded acceptable preservation parameters prior to pre-treatment (Table S3). During post-treatment, the majority of the analysed samples IRSF, C/C, and BPI parameters fell outside of the accepted ranges after the acetic acid (buffered and unbuffered, 15 min, 4 h, 8 h), NaClO, and combined treatments. This includes both modern and archaeological samples, suggesting that the changes are caused by alterations of the indigenous enamel structures and not just removal of contamination. In almost all samples (over 90%), IRSF values increased up to 0.7 after treatment. Samples that failed FTIR parameters prior to treatment generally showed no significant improvement or decline following pre-treatment.

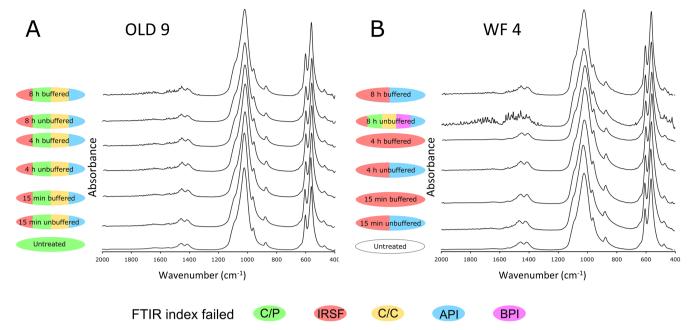


FIGURE 4 | Examples of ATR-FTIR spectra exhibiting various modes of diagenetic alteration for powdered enamel bioapatite samples subjected to different acetic acid treatments from (a) Neolithic Oldenburg LA 77 (OLD 9) and (b) PPNA WF16 (WF 4). Colours indicate failed FTIR preservation indices: C/P, IRSF, C/C, API, BPI.

5 | Discussion

5.1 | Pre-Treatments for Removing Organic Matter From Bioapatite

Organic removal protocols using $\rm H_2O_2$ or NaClO are regularly used to prepare tooth enamel samples for stable isotope analysis. However, our results, together with previously published studies, demonstrate that the use of oxidising reagents to remove organic contamination from enamel samples is not recommended. Tooth enamel contains a very small amount of organic content (<1%) compared to 20–25% in dentine and bone [18]. Considering that the organic content in teeth and bone is comprised largely of collagenous proteins that degrade in soils and sediments over time [93], the organic content present in archaeological enamel samples is negligible. Furthermore, previous work has demonstrated that only samples with a high organic content, where the total inorganic carbon to total organic carbon ratio exceeds 0.3, are likely to exhibit meaningful shifts in carbonate oxygen and carbon isotope ratios [94].

The use of an oxidising reagent step for organic removal during enamel sample preparation, however, may alter the chemical properties of the enamel. Here, samples treated with NaClO exhibited significantly less weight loss compared to those subjected to other treatment protocols, with an average mass loss of 18.3% compared to 22–23% documented for other treatments. The lower sample loss could be attributed to adsorption of exogenous carbon from the atmosphere [14, 15, 95]. The high pH of the NaClO solution promotes the dissolution of atmospheric CO_2 , which is subsequently converted into carbonic acid and then bicarbonate and carbonate ions, which can precipitate as secondary calcium carbonate (CaCO_3) onto the powdered sample [12]. This precipitation of secondary carbonates is further corroborated by previous studies reporting higher carbonate

content (%CO₃) in sodium hypochlorite-treated samples compared to untreated controls [15, 60, 95].

5.2 | Pre-Treatments to Remove Secondary and Exogenous Carbonates From Bioapatite

The results of this study demonstrate that the rate of sample dissolution is similar between 0.1 M acetic acid and 1 M calcium acetate buffered acetic acid treatments. Unbuffered acetic acid has a low pH value (~ 3) which may promote sample loss, particularly problematic for sequentially sampled teeth where sample yields are already low. Snoeck and Pellegrini [60] found that unbuffered acetic acid imparts a drastically larger sample weight loss, with an average enamel sample loss of 65% compared with a 30% sample loss associated with buffered acetic solution after only a 30 min treatment length. In contrast, our results show that there are no significant differences in the rate of dissolution in tooth enamel samples between buffered and unbuffered acetic acid solutions regardless of treatment length, with an average sample loss of 23% for both treatments. Given the absence of significant differences in sample weight loss across treatment durations, it is plausible that most dissolution occurs within the first 15 min of the reaction. This observation is consistent with findings by Garvie-Lok et al. [58] and Skippington et al. [13], who reported that enamel samples exhibit the highest reactivity during the first 15 min of treatment, with dissolution equilibrium reached after 4h of treatment.

The absence of a correlation between treatment duration and the magnitude of carbon and oxygen isotope change in samples treated with unbuffered acetic acid (r =0.08), along with only a weak negative correlation in buffered acetic acid-treated samples (r =-0.13), suggests that more than just diagenetic matter is being removed from the samples. The smaller Δ^{13} C

and Δ^{18} O values after 4h or 8h treatment durations compared to the 15 min treatment may indicate an initial dissolution of exogenous (diagenetic) carbonate followed by slower modification of structural bioapatite through dissolution or recrystallisation [13]. As diagenetic carbonates tend to be more soluble, non-lattice bound ions would dissolve first causing the initial change in bioapatite isotope composition [15]. Any subsequent modification of bioapatite, including dissolution of structural carbonate or recrystallisation, can affect carbon and oxygen isotope values, as reflected in Δ^{13} C and Δ^{18} O values observed after 4h and 8h of treatment [15, 56, 57]. It is important to note, however, that in some samples these shifts are negligible or very small. Given that the 1σ analytical uncertainty, calculated as the standard deviation of repeated measurements of both internal enamel standards and sample duplicates, is 0.1% for δ^{13} C and 0.2% for δ^{18} O, these observed small isotopic shifts are not considered significant.

Experimental data measured from modern tooth enamel samples demonstrate that δ^{13} C and δ^{18} O values of samples change after treatment with acetic acid solution [12, 13, 15, 56]. As modern tooth samples have not been subjected to potential contamination from a burial environment and have not undergone diagenesis, the isotopic changes observed in other studies and here cannot be assigned to the removal of exogenous carbonates. This demonstrates complex processes take place during pretreatment, including dissolution of exogenous carbonates and simultaneous reactions with structural carbonate (dissolution and/or recrystallisation) that affects the fidelity of the original in vivo isotope signal.

Changes in FTIR IRSF, C/C, and API indices following acetic acid treatment indicate alterations to the bioapatite structure caused by the treatment itself. An increase in IRSF, an indicator of crystal structural order, is expected with the removal of secondary, diagenetic, and non-lattice bound carbonate ions [14, 29, 83]. However, it is notable and concerning that several samples exhibiting pre-treatment IRSF values within the accepted range (3.1-3.9) exhibited elevated IRSF values posttreatment, ranging from 4.1 to 4.6. Although FTIR spectra of untreated and chemically treated bioapatite are not directly comparable due to differing preparation methods [83, 96], the magnitude of change observed in IRSF values suggests that chemical alteration beyond the removal of contaminants is occurring during treatment. Specifically, the increase in IRSF values is consistent with recrystallisation processes, which produce a more ordered crystalline structure. While our study did not observe brushite recrystallisation (see Lee-Thorp & van der Merwe [56]), the high IRSF (>4.0) values nonetheless indicate that some degree of recrystallisation is likely taking place during acetic acid exposure [97]. This warrants caution in interpreting isotopic data from treated samples, particularly when FTIR indices shift beyond established thresholds. When the supernatant from the acetic acid treatment is dried and analysed using ATR-FTIR, it becomes evident that structural components of the powdered enamel sample are being dissolved. The FTIR spectra of the supernatant closely resemble that of enamel bioapatite, exhibiting characteristic phosphate peaks at 1015, 565, and 605 cm⁻¹ as well as carbonate peaks at 1415, 1450, and 1550 cm⁻¹ (Figure S2). Notably, supernatant peak intensities, particularly the 1550 cm⁻¹ peak in carbonate region of the spectra, differ

significantly from those of raw enamel powder. This difference likely reflects the differential dissolution of type-A, OH- substituting carbonate associated with the 1450 and 1550 cm⁻¹ peaks versus type-B, phosphate-substituting carbonate associated with the 1415 cm⁻¹ peak. Type A carbonate ions may dissolve more readily due to their larger size and greater structural disorder relative to the OH- ions they replace [83]. Additionally, Amide I and Amide II bands, indicative of organic material, appear in the 1400–1700 cm⁻¹ range. While some spectral variation may be attributable to the removal of residual organics, the overall organic content in enamel bioapatite is minimal and does not sufficiently account for the pronounced 1550 cm⁻¹ peak. These findings suggest that structural components, including carbonate substitutions, are being partially dissolved during acetic acid treatment.

5.3 | Combined Organic and Secondary Mineral Pre-Treatment

When acetic acid is applied as a secondary treatment step following NaClO pre-treatment, the directionality in carbon isotopic change is the opposite of that observed for the NaClO-only pre-treatment. The combined treatment also results in the same weight loss comparable to that observed for the acetic acid-only treatment, whereas NaClO treatment alone yields a lower average weight loss of approximately 16%. The application of sodium hypochlorite to enamel samples induces dissolution of atmospheric CO₂, which subsequently precipitates as carbonate ions in the solution, it is therefore likely that the subsequent acetic acid treatment removes these precipitated carbonates [12, 60, 98]. Since these precipitated carbonates are not structurally bound to the enamel bioapatite, they would be dissolved along with any diagenetic carbonates. However, complete removal of all precipitated carbonates during this step cannot be guaranteed. Given the negligible organic content in tooth enamel and the potential for secondary carbonate precipitation, the application of NaClO, either alone or in combination with acetic acid, should be avoided in enamel sample preparation.

5.4 | Does Burial Environment Affect the Preservation of the Isotopic Signature?

FTIR indices indicate that, regardless of burial environment, archaeological tooth enamel samples in this study have been affected by diagenesis or other transformation processes. At THS, the presence of calcic [67] soils likely contributed to the precipitation of calcareous concretions on tooth enamel surfaces, which FTIR analysis identified as carbonated apatite (Figure S3). FTIR screening of the THS enamel powders also showed some diagenetic alteration in the enamel structure, possibly related to these concretions. The average δ^{13} C and δ^{18} O values of the concretions, -8.7% and -2.0%, respectively (Table S4), closely resemble those of the corresponding untreated enamel sample, suggesting possible contamination by carbonated apatite. Following acetic acid treatment, enamel δ^{13} C values were depleted in 13 C and enriched in ¹⁸O, which may indicate contamination originating from the concretions was removed. However, since δ^{13} C and δ^{18} O values of untreated enamel and adhering concretions were so similar, it is also possible that the isotope value change is due to the removal of structural carbonate during acetic acid treatment rather than removal of contamination. Supporting this interpretation, samples from WF16 recovered from a similarly carbonate-rich environment but lacking visible concretions or FTIR signatures indicating diagenetic alteration, also exhibited negative $\delta^{13}{\rm C}$ shifts of a similar magnitude and positive $\delta^{18}{\rm O}$ shifts of smaller magnitude as the THS samples. This pattern suggests that isotopic changes associated with chemical pre-treatment may primarily result from partial dissolution and/or recrystallisation processes induced by the pre-treatment itself, rather than removal of diagenetic carbonates.

The Oldenburg LA 77 samples appear to have undergone a distinctly different diagenetic pathway compared to those from THS and WF16, likely due to fluctuations in the local water table and the relatively low carbonate content of the burial environment sediments. The brown to black staining present on the teeth is consistent with contamination by iron and manganese oxides which is commonly found in waterlogged dental and osseous tissues [74–77]. Water movement through pore space in bioapatite structures can lead to the dissolution of endogenous enamel structure and the incorporation of exogenous compounds via ionic exchange processes [25]. Evidence for post-depositional alteration is further supplied by numerous FTIR indices in the Oldenburg samples falling outside of accepted reference ranges.

Overall, this study demonstrates that while FTIR pre-screening of enamel samples can reveal certain structural diagenetic alteration, it does not directly indicate modification to the in vivo carbon and oxygen isotope ratios. Furthermore, FTIR analysis has limited capacity to distinguish exogenous carbonate contamination, as soil-derived carbonates exhibit overlapping absorbance peaks, such as the characteristic $1415\,\mathrm{cm}^{-1}$ peak shared by enamel bioapatite and other carbonate minerals (CaCO₃, MgCO₃) [83].

6 | Conclusions

Based on findings from this study as well as previously published data sets, we observe:

- Pre-treatments meant to remove organic contamination, such as H₂O₂ and NaClO, should be avoided as they incur unwanted changes to the chemical composition and isotope values of tooth enamel carbonates.
- Pre-treatments involving the use of acetic acid, while likely efficient in removing secondary carbonates, impart some changes in enamel carbonate $\delta^{13}C$ and $\delta^{18}O$ values. Isotopic shifts observed in both modern and archaeological tooth samples suggest that, in addition to removing exogenous carbonates, acetic acid also dissolves portions of structural carbonate or induces other modifications such as recrystallisation or ionic exchange.
- Both buffered and unbuffered acetic acids are suitable solvents as they affect sample loss, structure and isotope values in a comparable manner.
- While no differences between acetic acid treatment lengths were found in this study, we, in agreement with previous

researchers, recommend shorter duration treatments (e.g., 15 min) to minimise the risk of structural enamel carbonate alteration, dissolution, and recrystallisation.

We also conclude that ATR-FTIR pre-screening, while useful in assessing the structural integrity of the sample and diagenetic alteration, might not directly establish the preservation of in vivo oxygen and carbon isotope ratios. Nonetheless, we strongly recommend ATR-FTIR pre-screening as a means to document and critically evaluate potential vectors of isotopic change. It is also clear that chemical pre-treatment-induced changes to δ^{13} C and δ^{18} O values are influenced by environmental variables and burial sedimentary context, the latter of which may vary across a single archaeological site. Therefore, careful consideration of site formation processes, sediment mineral composition, and perhaps also sediment isotopic composition is essential for understanding the diagenetic pathways affecting archaeological enamel. Finally, we emphasise the urgent need to standardise protocols for the preparation of enamel bioapatite samples to ensure reproducibility of results and minimise interpretative bias.

Author Contributions

Karolina Varkulevičiūtė: conceptualization, methodology, formal analysis, investigation, visualization, writing – original draft, writing – review and editing. **Christine Winter-Schuh:** conceptualization, methodology, data curation, investigation, writing – review and editing. **Cheryl A. Makarewicz:** conceptualization, methodology, data curation, investigation, supervision, writing – review and editing.

Acknowledgements

This study was funded by the Deutsche Forschungsgemeinschaft (DFG, German Research Foundation) under Germany's Excellence Strategy—EXC 2150-390870439. Open Access funding enabled and organized by Projekt DEAL.

Data Availability Statement

The data that support the findings of this study are available from the corresponding author upon reasonable request.

Peer Review

The peer review history for this article is available at https://www.webofscience.com/api/gateway/wos/peer-review/10.1002/rcm.10090.

References

- 1. A. Zazzo, M. Balasse, B. H. Passey, A. P. Moloney, F. J. Monahan, and O. Schmidt, "The Isotope Record of Short- and Long-Term Dietary Changes in Sheep Tooth Enamel: Implications for Quantitative Reconstruction of Paleodiets," *Geochimica et Cosmochimica Acta* 74, no. 12 (2010): 3571–3586, https://doi.org/10.1016/j.gca.2010.03.017.
- 2. J. E. Laffoon, R. V. Rojas, and C. L. Hofman, "Oxygen and Carbon Isotope Analysis of Human Dental Enamel from the Caribbean: Implications for Investigating Individual Origins," *Archaeometry* 55, no. 4 (2013): 742–765, https://doi.org/10.1111/j.1475-4754.2012.00698.x.
- 3. S. E. Pilaar Birch, P. T. Miracle, R. E. Stevens, and T. C. O'Connell, "Late Pleistocene/Early Holocene Migratory Behavior of Ungulates Using Isotopic Analysis of Tooth Enamel and Its Effects on Forager Mobility," *PLoS ONE* 11, no. 6 (2016): e0155714, https://doi.org/10.1371/journal.pone.0155714.

- 4. P. Heddell-Stevens, O. Jöris, K. Britton, et al., "Multi-Isotope Reconstruction of Late Pleistocene Large-Herbivore Biogeography and Mobility Patterns in Central Europe," *Communications Biology* 7, no. 1 (2024): 568, https://doi.org/10.1038/s42003-024-06233-2.
- 5. M. Pellegrini, J. Pouncett, M. Jay, M. P. Pearson, and M. P. Richards, "Tooth Enamel Oxygen "Isoscapes" Show a High Degree of Human Mobility in Prehistoric Britain," *Scientific Reports* 6, no. 1 (2016): 34986, https://doi.org/10.1038/srep34986.
- 6. L. Dai, M. Balasse, J. Yuan, C. Zhao, Y. Hu, and J. D. Vigne, "Cattle and Sheep Raising and Millet Growing in the Longshan Age in Central China: Stable Isotope Investigation at the Xinzhai Site," *Quaternary International* 426 (2016): 145–157, https://doi.org/10.1016/j.quaint.2016.
- 7. L. Jacques, N. Ogle, I. Moussa, et al., "Implications of Diagenesis for the Isotopic Analysis of Upper Miocene Large Mammalian Herbivore Tooth Enamel from Chad," *Palaeogeography, Palaeoclimatology, Palaeoceology* 266, no. 3–4 (2008): 200–210, https://doi.org/10.1016/j.palaeo. 2008.03.040.
- 8. J. Aufort, C. Gervais, L. Ségalen, et al., "Atomic Scale Transformation of Bone in Controlled Aqueous Alteration Experiments," *Palaeogeography, Palaeoclimatology, Palaeoecology* 526 (2019): 80–95, https://doi.org/10.1016/j.palaeo.2019.03.042.
- 9. C. N. Trueman, L. Kocsis, M. R. Palmer, and C. Dewdney, "Fractionation of Rare Earth Elements within Bone Mineral: A Natural Cation Exchange System," *Palaeogeography, Palaeoclimatology, Palaeocology* 310, no. 1–2 (2011): 124–132, https://doi.org/10.1016/j.palaeo.2011. 01.002.
- 10. G. Dal Sasso, M. Lebon, I. Angelini, L. Maritan, D. Usai, and G. Artioli, "Bone Diagenesis Variability Among Multiple Burial Phases at Al Khiday (Sudan) Investigated by ATR-FTIR Spectroscopy," *Palaeogeography, Palaeoclimatology, Palaeoecology* 463 (2016): 168–179, https://doi.org/10.1016/j.palaeo.2016.10.005.
- 11. M. Lebon, A. Zazzo, and I. Reiche, "Screening In Situ Bone and Teeth Preservation by ATR-FTIR Mapping," *Palaeogeography, Palaeoclimatology, Palaeoecology* 416 (2014): 110–119, https://doi.org/10.1016/j.palaeo.2014.08.001.
- 12. M. Pellegrini and C. Snoeck, "Comparing Bioapatite Carbonate Pre-Treatments for Isotopic Measurements: Part 2 Impact on Carbon and Oxygen Isotope Compositions," *Chemical Geology* 420 (2016): 88–96, https://doi.org/10.1016/j.chemgeo.2015.10.038.
- 13. J. Skippington, P. Veth, T. Manne, and M. Slack, "Preanalytical Processing of Archaeological Mammal Enamel Apatite Carbonates for Stable Isotope Investigations: A Comparative Analysis of the Effect of Acid Treatment on Samples from Northwest Australia," *International Journal of Osteoarchaeology* 29, no. 5 (2019): 760–771, https://doi.org/10.1002/oa.2787.
- 14. V. Grimes and M. Pellegrini, "A Comparison of Pretreatment Methods for the Analysis of Phosphate Oxygen Isotope Ratios in Bioapatite," *Rapid Communications in Mass Spectrometry* 27, no. 3 (2013): 375–390, https://doi.org/10.1002/rcm.6463.
- 15. B. E. Crowley and P. V. Wheatley, "To Bleach or Not to Bleach? Comparing Treatment Methods for Isolating Biogenic Carbonate," *Chemical Geology* 381 (2014): 234–242, https://doi.org/10.1016/j.chemgeo.2014. 05.006.
- 16. J. H. Bowes and M. M. Murray, "The Chemical Composition of Teeth: The Composition of Human Enamel and Dentine," *Biochemical Journal* 29, no. 12 (1935): 2721–2727, https://doi.org/10.1042/bj0292721.
- 17. M. J. Kohn, M. J. Schoeninger, and W. W. Barker, "Altered States: Effects of Diagenesis on Fossil Tooth Chemistry," *Geochimica et Cosmochimica Acta* 63, no. 18 (1999): 2737–2747, https://doi.org/10.1016/S0016-7037(99)00208-2.
- 18. Y. Wang and T. E. Cerling, "A Model of Fossil Tooth and Bone Diagenesis: Implications for Paleodiet Reconstruction From Stable

- Isotopes," *Palaeogeography, Palaeoclimatology, Palaeoecology* 107, no. 3–4 (1994): 281–289, https://doi.org/10.1016/0031-0182(94)90100-7.
- 19. P. D. Frazier, "Adult Human Enamel: An Electron Microscopy Study of Crystallite Size and Morphology," *Journal of Untrastructure Research* 22 (1968): 1–11.
- 20. W. V. Koenigswald and W. A. Clemens, "Levels of Complexity in the Microstructure of Mammalian Enamel and Their Application in Studies of Systematics," *Scanning Microscopy* 6, no. 1 (1992): 195–218.
- 21. W. J. Hamilton, Jr., G. Judd, and G. S. Ansell, "Ultrastructure of Human Enamel Specimens Prepared by Ion Micromilling," *Journal of Dental Research* 52, no. 4 (1973): 703–710, https://doi.org/10.1177/00220 345730520041101.
- 22. R. P. Shellis and G. H. Didbin, "Enamel Microporosity and Its Functional Implications," in *Development, Function and Evolution of Teeth*, eds. M. F. Teaford, M. M. Smith, and M. W. Ferguson (Cambridge University Press, 2000), 242–251.
- 23. G. Bergman and B. Siljestrand, "Water Evaporation in Vitro from Human Dental Enamel," *Archives of Oral Biology* 8 (1963): 37–38, https://doi.org/10.1016/0003-9969(63)90090-6.
- 24. D. F. G. Poole, P. W. Tailby, and D. C. Berry, "The Movement of Water and Other Molecules Through Human Enamel," *Archives of Oral Biology* 8, no. 6 (1963): 771–772, https://doi.org/10.1016/0003-9969(63) 90009-8.
- 25. R. E. M. Hedges and A. R. Millard, "Bones and Groundwater: Towards the Modelling of Diagenetic Processes," *Journal of Archaeological Science* 22, no. 2 (1995): 155–164, https://doi.org/10.1006/jasc.1995.0017.
- 26. H. W. Krueger, "Exchange of Carbon With Biological Apatite," *Journal of Archaeological Science* 18, no. 3 (1991): 355–361, https://doi.org/10.1016/0305-4403(91)90071-V.
- 27. B. K. Nelson, M. J. Deniro, M. J. Schoeninger, D. J. De Paolo, and P. E. Hare, "Effects of Diagenesis on Strontium, Carbon, Nitrogen and Oxygen Concentration and Isotopic Composition of Bone," *Geochimica et Cosmochimica Acta* 50, no. 9 (1986): 1941–1949, https://doi.org/10.1016/0016-7037(86)90250-4.
- 28. S. Decrée, D. Herwartz, J. Mercadier, et al., "The Post-Mortem History of a Bone Revealed by Its Trace Element Signature: The Case of a Fossil Whale Rostrum," *Chemical Geology* 477 (2018): 137–150, https://doi.org/10.1016/j.chemgeo.2017.12.021.
- 29. M. Sponheimer and J. A. Lee-Thorp, "Oxygen Isotopes in Enamel Carbonate and Their Ecological Significance," *Journal of Archaeological Science* 26, no. 6 (1999): 723–728, https://doi.org/10.1006/jasc. 1998.0388.
- 30. H. U. Pfretzschner, "Fossilization of Haversian Bone in Aquatic Environments," *Comptes Rendus Palevol* 3, no. 6–7 (2004): 605–616, https://doi.org/10.1016/j.crpv.2004.07.006.
- 31. R. Millard and E. M. Hedges, "A Diffusion-Adsorption Model of Uranium Uptake by Archaeological Bone," *Geochimica Et Cosmochimica Acta* 60, no. 12 (1996): 2139–2159, https://doi.org/10.1016/0016-7037(96) 00050-6.
- 32. M. J. Kohn, "Models of Diffusion-Limited Uptake of Trace Elements in Fossils and Rates of Fossilization," *Geochimica et Cosmochimica Acta* 72, no. 15 (2008): 3758–3770, https://doi.org/10.1016/j.gca.2008.05.045.
- 33. R. Madgwick, J. Mulville, and J. Evans, "Investigating Diagenesis and the Suitability of Porcine Enamel for Strontium (87Sr/86Sr) Isotope Analysis," *Journal of Analytical Atomic Spectrometry* 27, no. 5 (2012): 733, https://doi.org/10.1039/c2ja10356g.
- 34. G. D. Kamenov, E. M. Lofaro, G. Goad, and J. Krigbaum, "Trace Elements in Modern and Archaeological Human Teeth: Implications for Human Metal Exposure and Enamel Diagenetic Changes," *Journal of Archaeological Science* 99 (2018): 27–34, https://doi.org/10.1016/j.jas. 2018.09.002.

- 35. K. Weber, M. Weber, M. Menneken, et al., "Diagenetic Stability of Non-Traditional Stable Isotope Systems (Ca, Sr, Mg, Zn) in Teeth—An In-Vitro Alteration Experiment of Biogenic Apatite in Isotopically Enriched Tracer Solution," *Chemical Geology* 572 (2021): 120196, https://doi.org/10.1016/j.chemgeo.2021.120196.
- 36. A. Zazzo, "Bone and Enamel Carbonate Diagenesis: A Radiocarbon Prospective," *Palaeogeography, Palaeoclimatology, Palaeoecology* 416 (2014): 168–178, https://doi.org/10.1016/j.palaeo.2014.05.006.
- 37. Z. D. Sharp, "The Effect of Diagenesis on Oxygen Isotope Ratios of Biogenic Phosphates," *American Journal of Science* 300, no. 3 (2000): 222–237, https://doi.org/10.2475/ajs.300.3.222.
- 38. S. E. Etok, E. Valsami-Jones, T. J. Wess, et al., "Structural and Chemical Changes of Thermally Treated Bone Apatite," *Journal of Materials Science* 42, no. 23 (2007): 9807–9816, https://doi.org/10.1007/s1085 3-007-1993-z.
- 39. F. Berna, A. Matthews, and S. Weiner, "Solubilities of Bone Mineral From Archaeological Sites: The Recrystallization Window," *Journal of Archaeological Science* 31, no. 7 (2004): 867–882, https://doi.org/10.1016/j.jas.2003.12.003.
- 40. A. Zazzo, C. Lécuyer, and A. Mariotti, "Experimentally-Controlled Carbon and Oxygen Isotope Exchange Between Bioapatites and Water Under Inorganic and Microbially-Mediated Conditions," *Geochimica et Cosmochimica Acta* 68, no. 1 (2004): 1–12, https://doi.org/10.1016/S0016-7037(03)00278-3.
- 41. C. M. Nielsen-Marsch and R. E. M. Hedges, "Patterns of Diagenesis in Bone II: Effects of Acetic Acid Treatment and the Removal of Diagenetic CO₃²⁻," *Journal of Archaeological Science* 27, no. 12 (2000): 1151–1159.
- 42. L. E. Wright and H. P. Schwarcz, "Infrared and Isotopic Evidence for Diagenesis of Bone Apatite at Dos Pilas, Guatemala: Palaeodietary Implications," *Journal of Archaeological Science* 23, no. 6 (1996): 933–944, https://doi.org/10.1006/jasc.1996.0087.
- 43. T. J. U. Thompson, M. Islam, K. Piduru, and A. Marcel, "An Investigation into the Internal and External Variables Acting on Crystallinity Index Using Fourier Transform Infrared Spectroscopy on Unaltered and Burned Bone," *Palaeogeography, Palaeoclimatology, Palaeoecology* 299, no. 1–2 (2011): 168–174, https://doi.org/10.1016/j.palaeo.2010.10.044.
- 44. M. Greiner, A. Rodríguez-Navarro, M. F. Heinig, et al., "Bone Incineration: An Experimental Study on Mineral Structure, Colour and Crystalline State," *Journal of Archaeological Science: Reports* 25 (2019): 507–518, https://doi.org/10.1016/j.jasrep.2019.05.009.
- 45. C. N. G. Trueman, A. K. Behrensmeyer, N. Tuross, and S. Weiner, "Mineralogical and Compositional Changes in Bones Exposed on Soil Surfaces in Amboseli National Park, Kenya: Diagenetic Mechanisms and the Role of Sediment Pore Fluids," *Journal of Archaeological Science* 31, no. 6 (2004): 721–739, https://doi.org/10.1016/j.jas.2003.11.003.
- 46. L. E. Munro, F. J. Longstaffe, and C. D. White, "Burning and Boiling of Modern Deer Bone: Effects on Crystallinity and Oxygen Isotope Composition of Bioapatite Phosphate," *Palaeogeography, Palaeoclimatology, Palaeoecology* 249, no. 1–2 (2007): 90–102, https://doi.org/10.1016/j.palaeo.2007.01.011.
- 47. J. Reyes-Gasga, E. L. Martínez-Piñeiro, G. Rodríguez-Álvarez, G. E. Tiznado-Orozco, R. García-García, and E. F. Brès, "XRD and FTIR Crystallinity Indices in Sound Human Tooth Enamel and Synthetic Hydroxyapatite," *Materials Science & Engineering, C: Materials for Biological Applications* 33, no. 8 (2013): 4568–4574, https://doi.org/10.1016/j.msec.2013.07.014.
- 48. C. Tornero and M. Saňa, "Integrating Stable Isotopes to the Study of the Origin of Management Strategies for Domestic Animals: $\delta^{13} C$ and $\delta^{18} O$ Results from Bioapatite Enamel of Cattle from the Tell Halula Site, Syria (7800–7000 BC)," in Proceedings of the 37th International Symposium on Archaeometry, 13th 16th May 2008, Siena, Italy, ed. I.

- Turbanti-Memmi (Springer Berlin Heidelberg, 2011), 435–440, https://doi.org/10.1007/978-3-642-14678-7_63.
- 49. M. Balasse, G. Obein, J. Ughetto-Monfrin, and I. Mainland, "Investigating Seasonality and Season of Birth in Past Herds: A Reference Set of Sheep Enamel Stable Oxygen Isotope Ratios," *Archaeometry* 54, no. 2 (2012): 349–368, https://doi.org/10.1111/j.1475-4754.2011.00624.x.
- 50. R. J. Stark, M. V. Emery, H. Schwarcz, et al., "Dataset of Oxygen, Carbon, and Strontium Isotope Values from the Imperial Roman Site of Velia (ca. 1st-2nd c. CE), Italy," *Data in Brief* 38 (2021): 107421, https://doi.org/10.1016/j.dib.2021.107421.
- 51. B. H. Passey, T. E. Cerling, M. E. Perkins, M. R. Voorhies, J. M. Harris, and S. T. Tucker, "Environmental Change in the Great Plains: An Isotopic Record From Fossil Horses," *Journal of Geology* 110, no. 2 (2002): 123–140, https://doi.org/10.1086/338280.
- 52. B. H. Passey, T. F. Robinson, L. K. Ayliffe, et al., "Carbon Isotope Fractionation between Diet, Breath CO₂, and Bioapatite in Different Mammals," *Journal of Archaeological Science* 32, no. 10 (2005): 1459–1470, https://doi.org/10.1016/j.jas.2005.03.015.
- 53. R. S. Feranec and B. J. MacFadden, "Isotopic Discrimination of Resource Partitioning Among Ungulates in C_3 -Dominated Communities from the Miocene of Florida and California," *Paleobiology* 32, no. 2 (2006): 191–205, https://doi.org/10.1666/05006.1.
- 54. J. R. O'Neil, E. B. Reinhard, and R. E. Blake, "A Rapid and Precise Method of Oxygen Isotope Analysis of Biogenic Phosphate," *Israel Journal of Earth Sciences* 43, no. 3–4 (1994): 203–212.
- 55. R. P. Lopes, C. S. Scherer, J. C. Pereira, and S. R. Dillenburg, "Paleo-environmental Changes in the Brazilian Pampa Based on Carbon and Oxygen Stable Isotope Analysis of Pleistocene Camelid Tooth Enamel," *Journal of Quaternary Science* 38, no. 5 (2023): 702–718, https://doi.org/10.1002/jqs.3502.
- 56. J. A. Lee-Thorp and N. J. Van Der Merwe, "Aspects of the Chemistry of Modern and Fossil Biological Apatites," *Journal of Archaeological Science* 18, no. 3 (1991): 343–354, https://doi.org/10.1016/0305-4403(91) 90070-6.
- 57. P. L. Koch, N. Tuross, and M. L. Fogel, "The Effects of Sample Treatment and Diagenesis on the Isotopic Integrity of Carbonate in Biogenic Hydroxylapatite," *Journal of Archaeological Science* 24, no. 5 (1997): 417–429, https://doi.org/10.1006/jasc.1996.0126.
- 58. S. J. Garvie-Lok, T. L. Varney, and M. A. Katzenberg, "Preparation of Bone Carbonate for Stable Isotope Analysis: The Effects of Treatment Time and Acid Concentration," *Journal of Archaeological Science* 31, no. 6 (2004): 763–776, https://doi.org/10.1016/j.jas.2003.10.014.
- 59. C. J. Yoder and E. J. Bartelink, "Effects of Different Sample Preparation Methods on Stable Carbon and Oxygen Isotope Values of Bone Apatite: A Comparison of Two Treatment Protocols," *Archaeometry* 52, no. 1 (2010): 115–130, https://doi.org/10.1111/j.1475-4754.2009.00473.x.
- 60. C. Snoeck and M. Pellegrini, "Comparing Bioapatite Carbonate Pretreatments for Isotopic Measurements: Part 1—Impact on Structure and Chemical Composition," *Chemical Geology* 417 (2015): 394–403, https://doi.org/10.1016/j.chemgeo.2015.10.004.
- 61. R. Z. LeGeros, O. R. Trautz, E. Klein, and J. P. LeGeros, "Two Types of Carbonate Substitution in the Apatite Structure," *Experientia* 25, no. 1 (1969): 5–7, https://doi.org/10.1007/BF01903856.
- 62. T. E. Cerling, J. M. Harris, S. H. Ambrose, M. G. Leakey, and N. Solounias, "Dietary and Environmental Reconstruction with Stable Isotope Analyses of Herbivore Tooth Enamel from the Miocene Locality of Fort Ternan, Kenya," *Journal of Human Evolution* 33, no. 6 (1997): 635–650, https://doi.org/10.1006/jhev.1997.0151.
- 63. A. Zazzo, H. Bocherens, D. Billiou, et al., "Herbivore Paleodiet and Paleoenvironmental Changes in Chad During the Pliocene Using Stable Isotope Ratios of Tooth Enamel Carbonate," *Paleobiology* 26, no. 2

- (2000): 294–309, https://doi.org/10.1666/0094-8373(2000)026<0294: HPAPCI>2.0.CO:2.
- 64. M. S. Chesson, "Preliminary Results of Excavations at Tell el-Handaquq South (1993–96)," *Palestine Exploration Quarterly* 130, no. 1 (1998): 20–34, https://doi.org/10.1179/peq.1998.130.1.20.
- 65. M. Bar-Matthews and A. Ayalon, "Mid-Holocene Climate Variations Revealed by High-Resolution Speleothem Records From Soreq Cave, Israel and Their Correlation With Cultural Changes," *Holocene* 21, no. 1 (2011): 163–171, https://doi.org/10.1177/0959683610384165.
- 66. S. Riehl, R. Bryson, and K. Pustovoytov, "Changing Growing Conditions for Crops During the Near Eastern Bronze Age (3000–1200 BC): The Stable Carbon Isotope Evidence," *Journal of Archaeological Science* 35, no. 4 (2008): 1011–1022, https://doi.org/10.1016/j.jas.2007.07.003.
- 67. S. A. Khresat, "Calcic Horizon Distribution and Soil Classification in Selected Soils of North-Western Jordan," *Journal of Arid Environments* 47, no. 2 (2001): 145–152, https://doi.org/10.1006/jare.2000.0712.
- 68. M. S. Chesson, "Urban Households in Early Bronze Age Communities of Syro-Palestine" (PhD Dissertation, Harvard University, 1997).
- 69. J. P. Brozio, W. Dörfler, I. Feeser, W. Kirleis, S. Klooß, and J. Müller, "A Middle Neolithic Well From Northern Germany: A Precise Source to Reconstruct Water Supply Management, Subsistence Economy, and Deposition Practices," *Journal of Archaeological Science* 51 (2014): 135–153, https://doi.org/10.1016/j.jas.2013.03.029.
- 70. O. Jakobsen, "Die Grube-Wesseker Niederung (Oldenburger Graben, Ostholstein): Quartärgeologische Und Geoarchäologische Untersuchungen Zur Landschaftsgeschichte Vor Dem Hintergrund Des Anhaltenden Postglazialen Meeresspiegelanstiegs". PhD Dissertation, Kiel: Kiel University; 2004.
- 71. J. P. Brozio, Megalithanlagen und Siedlungsmuster im trichterbecherzeitlichen Ostholstein, Frühe Monumentalität und soziale Differenzierung Bd. 9 (Habelt, 2016).
- 72. M. Böse, J. Ehlers, and F. Lehmkuhl, "Holozäne Klima- und Landschaftsgeschichte," in *Deutschlands Norden von Erdaltertum Zur Geganwart* (Springer, 2016), 193–239.
- 73. K. Barber, F. Chambers, and D. Maddy, "Late Holocene Climatic History of Northern Germany and Denmark: Peat Macrofossil Investigations at Dosenmoor, Schleswig-Holstein, and Svanemose, Jutland," *Boreas* 33, no. 2 (2004): 132–144, https://doi.org/10.1080/0300948041 0001082.
- 74. T. L. Dupras and J. A. Schultz, eds., "Taphonomic Bone Staining and Color Changes in Forensic Contexts," in *Manual of Forensic Taphonomy*, 1st ed. (CRC Press, 2013), 330–355, https://doi.org/10.1201/b15424-16.
- 75. E. M. S. Beyer-Olsen and S. Risnes, "Occurrence and Distribution of Blackish Staining on the Crowns of Human Teeth Obtained from an Archeological Excavation of a Medieval Site in Norway," *European Journal of Oral Sciences* 101, no. 2 (1993): 65–71, https://doi.org/10.1111/j. 1600-0722.1993.tb01090.x.
- 76. E. M. Stermer, S. Risnes, and P. M. Fischer, "Trace Element Analysis of Blackish Staining on the Crowns of Human Archaeological Teeth," *European Journal of Oral Sciences* 104, no. 3 (1996): 253–261, https://doi.org/10.1111/j.1600-0722.1996.tb00075.x.
- 77. E. L. Brown, R. A. Dixon, and J. W. Birkett, "The Discolouration of Human Teeth from Archaeological Contexts: Elemental Analysis of a Black Tooth from a Roman Cranium Recovered From the River Witham, Lincoln, UK," *Journal of Anthropology* 2014 (2014): 1–7, https://doi.org/10.1155/2014/859153.
- 78. B. Finlayson, S. Mithen, M. Al-Najjar, S. Smith, and E. Jenkins, "The Origins, Development and Practice of Economic and Social Strategies in the Middle East From Earliest Times to the Modern Day Excavations at Wadi Faynan 16, a Pre-Pottery Neolithic a Site in Southern Jordan," *Bulletin for the Council for British Research in the Levant* 5, no. 1 (2010): 46–66, https://doi.org/10.1179/175272710X12828116506116.

- 79. K. Wicks, "The Radiocarbon Dates," in WF16. The Excavation of an Early Neolithin Settlement in Wadi Faynan, Southern Jordan: Stratigraphy, Chronology, Architecture and Burials (Council for British research in Levant, 2018), https://www.jstor.org/stable/j.ctv37vwf8b.45.
- 80. S. Smith, A. Wade, E. Black, D. Brayshaw, C. Rambeau, and S. Mithen, "From Global Climate Change to Local Impact in Wadi Faynan, Southern Jordan: Ten Millennia of Human Settlement in Its Hydrological Context," in *Water, Life and Civilisation*, 1st ed., eds. S. Mithen and E. Black (Cambridge University Press, 2011), 218–244, https://doi.org/10.1017/CBO9780511975219.015.
- 81. S. Allcock and E. S. He Sediments, WF16. The Excavation of an Early Neolithic Settlement in Wadi Faynan, Southern Jordan:Stratigraphy, Chronology, Architecture and Burials (Council for British research in Levant, 2018), https://www.jstor.org/stable/j.ctv37vwf8b.45.
- 82. I. Angerer, G. Han, I. Fujisaki, and K. Havstad, "Climate Change and Ecosystems of Asia with Emphasis on Inner Mongolia and Mongolia," *Rangelands* 30, no. 3 (2008): 46–51.
- 83. C. A. M. France, N. Sugiyama, and E. Aguayo, "Establishing a Preservation Index for Bone, Dentin, and Enamel Bioapatite Mineral Using ATR-FTIR," *Journal of Archaeological Science: Reports* 33 (2020): 102551, https://doi.org/10.1016/j.jasrep.2020.102551.
- 84. R. E. M. Hedges, J. Lee-Thorp, and N. C. Tuross, "Is Tooth Enamel Carbonate a Suitable Material for Radiocarbon Dating?," *Radiocarbon* 37, no. 2 (1995): 285–290, https://doi.org/10.1016/j.palaeo.2019.03.042.
- 85. J. Sealy, N. Naidoo, V. J. Hare, S. Brunton, and J. T. Faith, "Climate and Ecology of the Palaeo-Agulhas Plain from Stable Carbon and Oxygen Isotopes in Bovid Tooth Enamel From Nelson Bay Cave, South Africa," *Quaternary Science Reviews* 235 (2020): 105974, https://doi.org/10.1016/j.quascirev.2019.105974.
- 86. M. Balasse, S. H. Ambrose, A. B. Smith, and T. D. Price, "The Seasonal Mobility Model for Prehistoric Herders in the South-western Cape of South Africa Assessed by Isotopic Analysis of Sheep Tooth Enamel," *Journal of Archaeological Science* 29, no. 9 (2002): 917–932, https://doi.org/10.1006/jasc.2001.0787.
- 87. A. Demény, A. D. Gugora, D. Kesjár, C. Lécuyer, and F. Fourel, "Stable Isotope Analyses of the Carbonate Component of Bones and Teeth: The Need for Method Standardization," *Journal of Archaeological Science* 109 (2019): 104979, https://doi.org/10.1016/j.jas.2019.104979.
- 88. P. Iacumin, H. Bocherens, A. Mariotti, and A. Longinelli, "An Isotopic Palaeoenvironmental Study of Human Skeletal Remains From the Nile Valley," *Palaeogeography, Palaeoclimatology, Palaeoecology* 126, no. 1–2 (1996): 15–30, https://doi.org/10.1016/S0031-0182(96)00067-3.
- 89. J. Towers, M. Jay, I. Mainland, O. Nehlich, and J. Montgomery, "A Calf for All Seasons? The Potential of Stable Isotope Analysis to Investigate Prehistoric Husbandry Practices," *Journal of Archaeological Science* 38, no. 8 (2011): 1858–1868, https://doi.org/10.1016/j.jas.2011. 03.030.
- 90. J. Sealy, J. Lee-Thorp, E. Loftus, J. T. Faith, and C. W. Marean, "Late Quaternary Environmental Change in the Southern Cape, South Africa, from Stable Carbon and Oxygen Isotopes in Faunal Tooth Enamel from Boomplaas Cave," *Journal of Quaternary Science* 31, no. 8 (2016): 919–927, https://doi.org/10.1002/jqs.2916.
- 91. D. Feng, T. Tütken, N. Löffler, G. Tröster, and A. Pack, "Isotopically Anomalous Metabolic Oxygen in Marine Vertebrates as Physiology and Atmospheric Proxy," *Geochimica et Cosmochimica Acta* 328 (2022): 85–102, https://doi.org/10.1016/j.gca.2022.05.008.
- 92. D. Feng, J. Surma, T. Tütken, et al., "Triple Oxygen Isotopes of Modern Terrestrial Mammalian Tooth Enamel—New Implications for Paleoenvironmental and Physiological Research," *Geochimica et Cosmochimica Acta* 365 (2024): 21–34, https://doi.org/10.1016/j.gca.2023.11.025.
- 93. G. Turner-Walker, "The Chemical and Microbial Degradation of Bones and Teeth," in *Advances in Human Palaeopathology*, 1st ed., eds.

- R. Pinhasi and S. Mays (Wiley, 2007), 3–29, https://doi.org/10.1002/9780470724187.ch1.
- 94. M. Oehlerich, M. Baumer, A. Lücke, and C. Mayr, "Effects of Organic Matter on Carbonate Stable Isotope Ratios (δ^{13} C, δ^{18} O Values) Implications for Analyses of Bulk Sediments," *Rapid Communications in Mass Spectrometry* 27, no. 6 (2013): 707–712, https://doi.org/10.1002/rcm.6492.
- 95. A. Zazzo, M. Balasse, and W. P. Patterson, "The Reconstruction of Mammal Individual History: Refining High-Resolution Isotope Record in Bovine Tooth Dentine," *Journal of Archaeological Science* 33, no. 8 (2006): 1177–1187, https://doi.org/10.1016/j.jas.2005.12.006.
- 96. R. Wood, A. B. C. Fleury, S. Fallon, T. M. H. Nguyen, and A. T. Nguyen, "Do Weak or Strong Acids Remove Carbonate Contamination from Ancient Tooth Enamel More Effectively? The Effect of Acid Pretreatment of Radiocarbon and δ^{13} C Analyses," *Radiocarbon* 63, no. 3 (2021): 935–952, https://doi.org/10.1017/RDC.2021.32.
- 97. R. J. A. Hopkins, C. Snoeck, and T. F. G. Higham, "When Dental Enamel is Put to the Acid Test: Pretreatment Effects and Radiocarbon Dating," *Radiocarbon* 58, no. 4 (2016): 893–904, https://doi.org/10.1017/RDC.2016.52.
- 98. F. Spencer, K. Verostick, A. Serna, C. Stantis, and G. J. Bowen, "Effects of Particle Size, Storage Conditions, and Chemical Pretreatments on Carbon and Oxygen Isotopic Measurements of Modern Tooth Enamel," *Science & Justice* 64, no. 2 (2024): 193–201, https://doi.org/10.1016/j.scijus.2024.01.004.

Supporting Information

Additional supporting information can be found online in the Supporting Information section.